

CO₂ Sequestration Influence on Low Permeable Geological Formations

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Abstract—Carbon dioxide (CO₂) is a major greenhouse effective gas that contributes to Earth's global warming. The idea is to inject the CO₂ in to the low permeable geological formation to allow the sufficient time for the trapping mechanisms by controlling the migration of CO₂ towards surface. MATLAB Reservoir Simulation Toolbox (MRST) is used to study the trapping of CO₂ in geological formation and migration. MRST-co₂lab is a collection of open source computational tools for modelling large-scale and long-time migration of CO₂ in conductive aquifers, combining ideas from basin modelling, computational geometry, hydrology, and reservoir simulation. All computational methods implemented in MRST-co₂lab are formulated based on a hybrid 2D grid that represents the 3D aquifer in terms of its top and bottom surfaces. The main objective of this work is to identify the geologically potential storage spill point for the storing CO₂ and the migration. We considered a volume of 5km x 5km x 50 m having homogeneous porosity of 0.25 that lies below the 1000m depth from the surface. The top surface will be placed on the grid structure.

Keywords: Storage, migration, MRST-co₂lab, spill point, spill region.

1. INTRODUCTION

Activities such as burning of fossil fuels along with cement making have led to a significant increase in the CO₂ concentration in the atmosphere since the start of the industrial revolution [IPCC 2007]. CO₂ being a strong and long-lived Green House Gas (GHG), its increased concentrations in the atmosphere has led to increase in the global temperature [IPCC 2007]. Carbon Capture, Storage and Sequestration are presently viewed as an important strategy to reduce the concentration of CO₂ in the atmosphere. One way to conflict climate change is to prevent the release of CO₂ to the environment by storing it in natural underground reservoirs. This paper describes the concept and outlines some of the issues involved in estimating the trapping of CO₂ in geological formation. CO₂ emitted by large sources, such as power plants and other industrial processes, could be captured and stored underground using Depleted oil as well as gas fields, Coal seams and deep saline reservoirs. Shaping the maximum amount of CO₂ that can be injected and firmly contained is a key matter. When injected into the geological formations,

density differences will drive CO₂ to form a separate (the CO₂ plume) mobile phase, which is driven upwards by buoyancy. To prevent CO₂ from moving straight forwardly upward, it is injected into a low permeable formation bounded upwards by a sealing caprock. Below this seal, CO₂ will spread out and slowly migrate in the upslope direction. Disregarding other trapping mechanisms, this migration continues until the plume encounters a trap in the top surface where CO₂ will accumulate. Once a trap is filled, excess CO₂ will spill over and keep migrating upwards to the next trap, and so on until the top of the formation is reached. In the short term, structural and stratigraphic trapping are key mechanisms for geological storage of CO₂. Herein, our primary concern is to develop methods that can quickly compute bounds on the overall capacity for structural trapping and select good positions for placing injection points. These methods have later been combined with simulation tools based on an assumption of vertical equilibrium simulation [H. M. Nilsen et al] to provide flow visualization and strategy to select injection point, observe the plume behaviour of co₂ and provide the result of trapping.

2. GEOLOGICAL STORAGE OF CARBON DIOXIDE

One way to combat climate the change is to prevent the release of CO₂ to the atmosphere by storing it in natural underground reservoirs and catch the topographical stockpiling of CO₂ give an approach to withdraw from transmitting CO₂ into the climate [Capture of CO₂ Kelly *et al*] [Geological storage of CO₂ by IEA]. This segment looks at these procedures and their impact on land stockpiling of CO₂. Thusly, the experience picked up from existing profound liquid infusion activities is important as far as the style of operation and is of a comparative size to what might be required for geographical stockpiling of CO₂.

2.1. Physical Properties of CO₂

The physical state of CO₂ shifts with temperature and pressure. At ambient conditions CO₂ is a gas; however it becomes liquid at greater depth. At high temperature, CO₂ is a

supercritical fluid when pressure is high enough. The transition from one state to another relies on upon the geothermal slope. In most sequestration scenarios, CO₂ are infused in liquid structure (low T, humble to high P), yet it changes into a supercritical fluid as it is injected and warms to the temperature of the development. In saline aquifers and oil reservoirs, CO₂ is less dense than the in situ fluids, so it ascends to the base of the seal. Clearly, maintaining an impermeable caprock is essential to containing the buoyant CO₂.

2.2. CO₂ Migration Behaviour

When CO₂ is injected into deep geological formations, it displaces the pore fluid. Depending on the fluid's properties, CO₂ is either miscible that is it can mix completely to form a single liquid phase or immiscible, so the phases stay separate. At conditions expected for sequestration, CO₂ and water are immiscible [CO₂ Sequestration in Deep Sedimentary Formations Sally by M. Benson *et al*]. Oil and CO₂ may or may not be miscible, contingent upon the structure of the oil and the formation pressure. CO₂ and natural gas are miscible. When the fluids are miscible, the CO₂ eventually displaces nearly all of the original fluid. Injection of an immiscible liquid sidesteps some part of the pore space, catching some of the original fluid. With the limited exception of dry-gas supplies, most sequestration ventures will require immiscible displacement to one degree or another [Carbon Dioxide Sequestration by David R. Cole *et al*]. For example, although oil and CO₂ are miscible, the water that is almost always present in formations is not miscible with oil or CO₂ oil mixtures. Equilibration of CO₂ between oil and water depends on the composition of the oil.

3. TRAPPING MECHANISMS OF CO₂

These trapping processes take place over a lot of years at dissimilar rates from days to years to thousands of years, but in general, geologically stored CO₂ become more strongly trapped with time. Demonstration of a variety of geological storage of CO₂ is already being carried out in a range of projects of varying scale.

3.1. Structural/Stratigraphic trapping

Firstly, physical trapping of CO₂ below low permeability seals (cap rocks), such as very low permeability shale or salt beds, is the prime means to accumulate CO₂ in geological formations. Sedimentary basins have such closed, physically bound traps or structures, which are occupied mostly by saline water, oil and gas. Structural traps are those shaped by fold or fractured rocks. Faults can operate as permeability barriers in some circumstances and as preferential pathways for fluid flow in other conditions. Stratigraphic traps are formed by lateral changes in rock type caused by variation in the setting where the rocks were settled down. Both of these types of traps are good for CO₂ storage. A special folder for structural trapping can take place in saline formations that do not have a

closed trap but consists of a slightly tilted aquifer where fluids move very slowly over long distances. When CO₂ is injected into a formation, it dislocates saline formation water and subsequently migrates buoyantly upwards, since it is less dense than the water. When it reaches the top of the storage formation, it continue to migrate as a separate phase until it is dissolved (potentially helped by the gravity instability and mixing), trapped as residual CO₂ saturation or gets blocked in local stratigraphic traps under the sealing formation [IPCC 2005].

3.2. Residual trapping

Consider a medium that is primarily packed with water or brine. The solid grains are usually made of minerals that are obviously wetting to water and, consequently the medium is particularly water wet. During CO₂ injection into the aquifer or reservoir, the non wetting CO₂ phase occupies the pore space. This is a drainage process in which the only mechanism for displacement of water by CO₂ is piston type displacement the CO₂ invades the porous medium in the form of a uninterrupted, associated cluster. Water, still, remain not only in small pores that have not been packed with CO₂ but also in the corners and gaps of the pores that have been invaded. Consider now the dislocation of the CO₂ by water. During this process, there are some physical mechanisms by which the water can displace the CO₂ [Lenormand *et al.*, 1983]. In accumulation to piston type displacement, accommodating pore body filling and snap-off may take place. For water wet rocks, snap off is the leading mechanism [Al-Futaisi and Patzek, 2003; Valvatne and Blunt, 2004].

3.3. Solubility Trapping

The dissolution of CO₂ and other flue gas impurities into the pore water can lead to trapping by solubility. The amount of gas that can dissolve into the water depends on several factors, most particularly pressure, temperature, and salinity of the brine [e.g. Spycher *et al.* 2003; Lagneau *et al.* 2005; Koschel *et al.* 2006; Oldenburg 2007]. At the conditions expected for maximum geological sequestration, CO₂ solubility increases with rising pressure (i.e. depth) but decreases with rising temperature and salinity. Benchscale experiments reveal that CO₂ dissolution is quick at high pressure when the water and CO₂ shares the same pore space [Czernichowski-Lauriol *et al* (1996)]. Still, in a real injection system, CO₂ dissolution may be rate-limited by the magnitude of the contact area between the CO₂ and the fluid phase. The principal benefit of solubility trapping is that once the CO₂ is dissolved, there is less CO₂ subject to the buoyant forces that drive it upwards.

3.4. Mineral Trapping

This mechanism occurs when dissolved CO₂ reacts directly or indirectly with minerals in the geologic formation, helpful for the precipitation of carbonate minerals [Oelkers *et al.* 2008]. Mineral trapping is attractive because it could arrest CO₂ for very long periods [Gunter *et al.* 1997]. However, the process is

thought to be comparatively slow because it depends on dissolution of silicate minerals, so the global impact may not be realized for tens to hundreds of years or longer.

4. METHODOLOGY

Trapping can be estimated using simple geometrical algorithms that compute traps and catchment areas for the top surface of a grid model. Simple percolation methods can be used to estimate spill paths, assuming that CO₂ is injected at an infinitesimal rate. These methods are implemented in free software in the MRST [MRST and s.krogstad et al (2007)]. The MRST-CO₂lab module [SINTEF ICT (oct 2014)] includes a graphical and numerical interface for interactively exhibiting structural trapping, input routines for standard input formats, and scripts to be download, unpack and process publicly available datasets. [E. K. Halland (2011), G. Eigestad (2008)]. Altogether, MRST provides a wonderful platform for supporting research, which we demonstrate here in providing complete scripts for all realistic examples.

4.1. The Spill Point Approach to Trap Analysis

We use the word spill path to refer to the path follow by CO₂ under the caprock on its buoyant migration, assume infinitesimal flow. While a trap has been totally filled by CO₂, any other quantity entering the trap will direct to an equivalent amount exit. For infinitesimal quantity, the flow will go after a spill path out of the trap that either terminate in a different trap or exits the area. Like such, individual traps are able to be seen to be linked by spill paths, greatly the way lakes are associated by rivers. In line by means of the principle of industry standard reservoir descriptions. We refer to this mesh as a top surface grid when resulting from a 3D reservoir model [Halvor Møll Nilsen et al]. Discovering all traps and spill paths for a known model can be finished by inspecting the geometry of this mesh. Even though the process is simple in principle, it is responsive to small change in input information. The algorithms make like results in the majority cases, but now and then the differences are significant. The figure1 explains the selection of spill point and spill path which will allow more amount of CO₂ as the approach explained above.

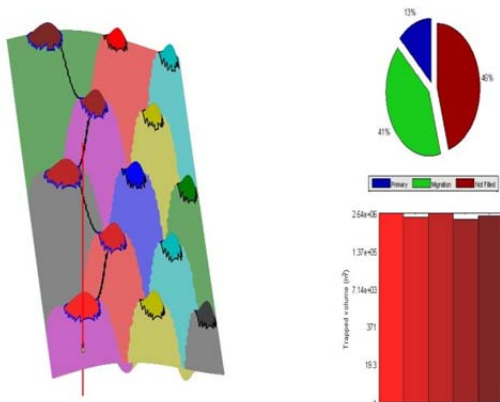


Fig. 1: Interactive vision in onward mode.

The pie chart presents the volume of the trap whose catchment area contains the injection point ('Primary'), the united volume of traps by the side of the associated spill path ('Migration'), and the volume of all other traps ('Not filled'). The bar plot shows the volume of each trap along the spill path. The scale is logarithmic because realistic scenario tends to have widely varying trap sizes.

4.2. Interface and implementation in MRST CO₂Lab

To utilize the functionality in MRST-CO₂lab, the first item one has to do is to make a semi-2D report of the top surface which includes a deposit of data objects that offer mapping between each cell in the 2D surface grid and a version of the volumetric column that lie under in 3D. The top surface grid can moreover be generate from a 3D volumetric grid by means of the function top Surface Grid. As a descriptive model of the type of information that can be obtain from this purpose. We consider box geometry of 5,000 × 5,000 × 50 m3. We initiate a sinusoidal perturbation of the top as well as bottom surfaces, shift the box to a depth of 1000 m also assign a uniform porosity of 0.25. Fig. 2 showsthe 3D grid along with the extracted top surface plotted on the similar axes. Assuming that the surface represents impermeable seal (Caprock), the local domes will symbolize structural traps.

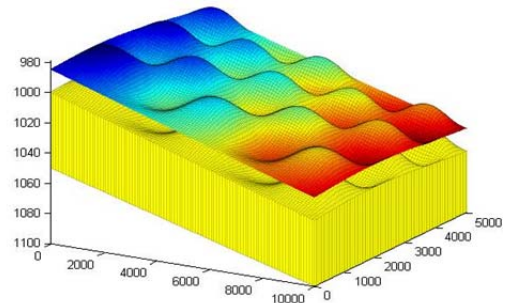


Fig. 2: Visionary model illuminate the trapping arrangement. The visualized extracted top-surface grid has been shifted slightly upwards and coloured by depth values to clearly discriminate it from the yellow 3D grid.

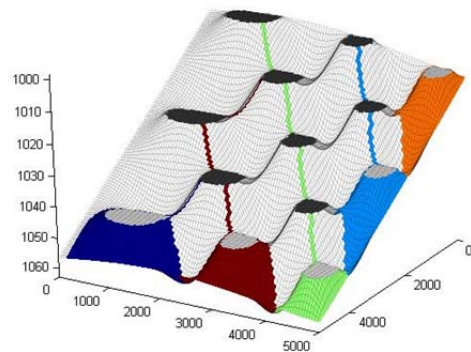


Fig. 3: Leaf traps depressed at the spill point (light gray), the neighbouring catchment areas (Shown in different colours), and the spill paths to the top of the field. The spill levels of up slope traps are exposed in dark gray.

The complete setup and all statements necessary to produce Figures 1, 2 and 3 can be found in the script trappingExample1.m. in MRST

5. RESULTS AND DISCUSSION

Through our simulations we have tried to express the flow of CO₂ in reservoir using vertical equilibrium model. The results were given for homogeneous aquifer in view of assuming the reservoir with geometric volume of $2.302e^{+09}$ m³ and 10000 cells were taken for constructing the reservoir. The injection site was located at cell: 1898 with(x: 9750, y: 925,

$i=98, j=19$). The rate of CO₂ injected per year is 0.3 megatonnes, pressure is 200 bars, migration time is 75 years and the injection time is 5 years. As it is a homogeneous and low permeable reservoir the permeability and porosity are maintained equal in all the cells permeability is taken 250mD (milli Darcy) and the porosity is taken as 0.25. the total trapping capacity is taken $4.948e^{+06}$ and this amounts to 39.58% of total pore volume of $1.25e^{+07}$. With these circumstances we have simulated the vertical equilibrium model and found the flow behaviour and type of trapping as shown in the Fig. 4 and Fig. 5 they describe behaviour of CO₂ at injection and migration period.

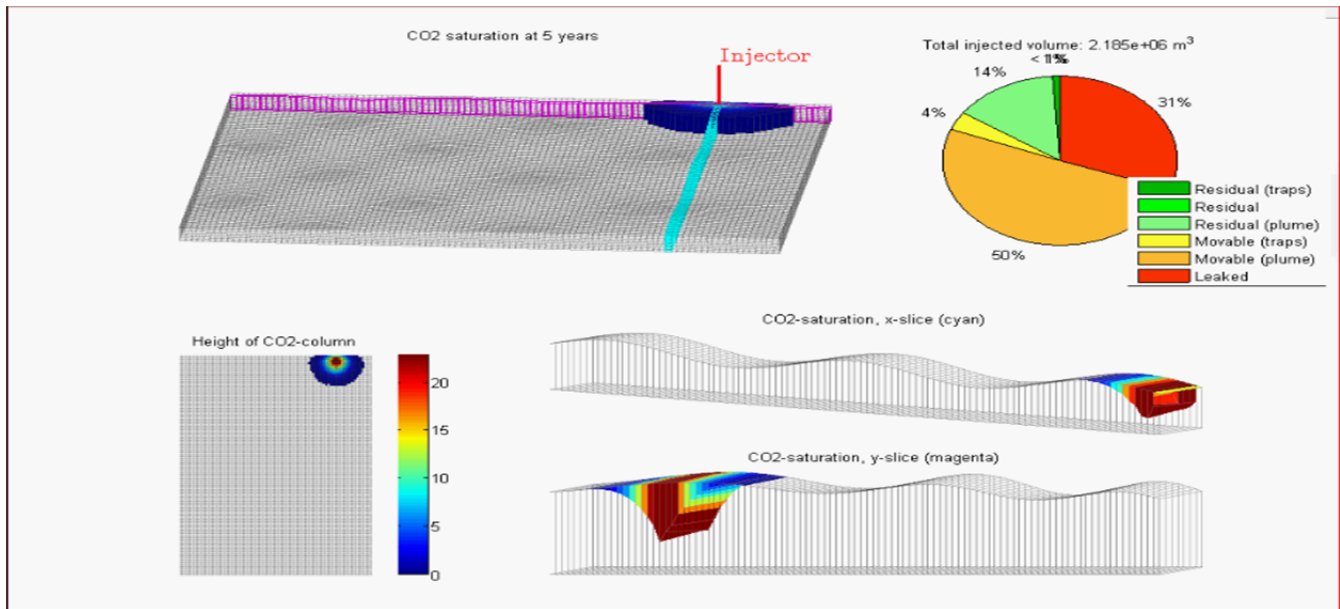


Fig. 4: The result of simulation at the end of injection. It illustrates the height of CO₂ at injection point is more and due to buoyant effect the CO₂ is flowing towards the cap rock

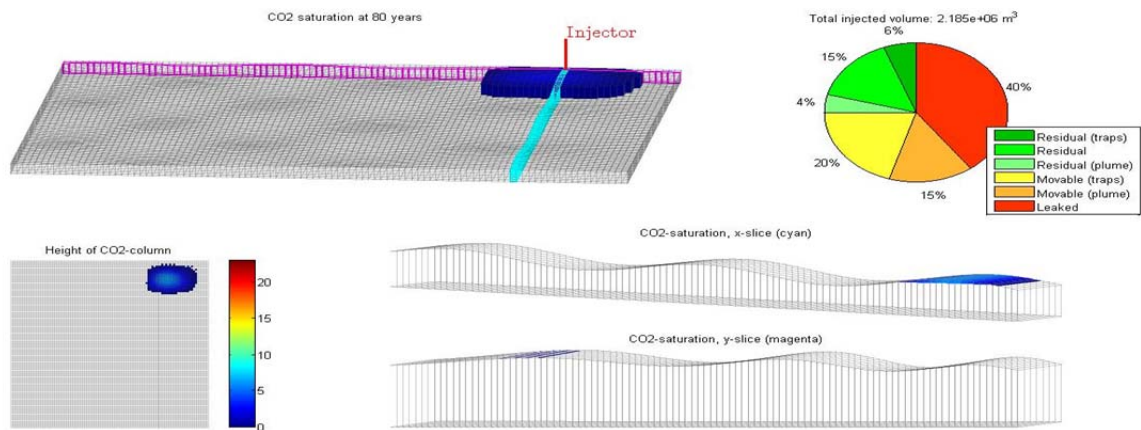
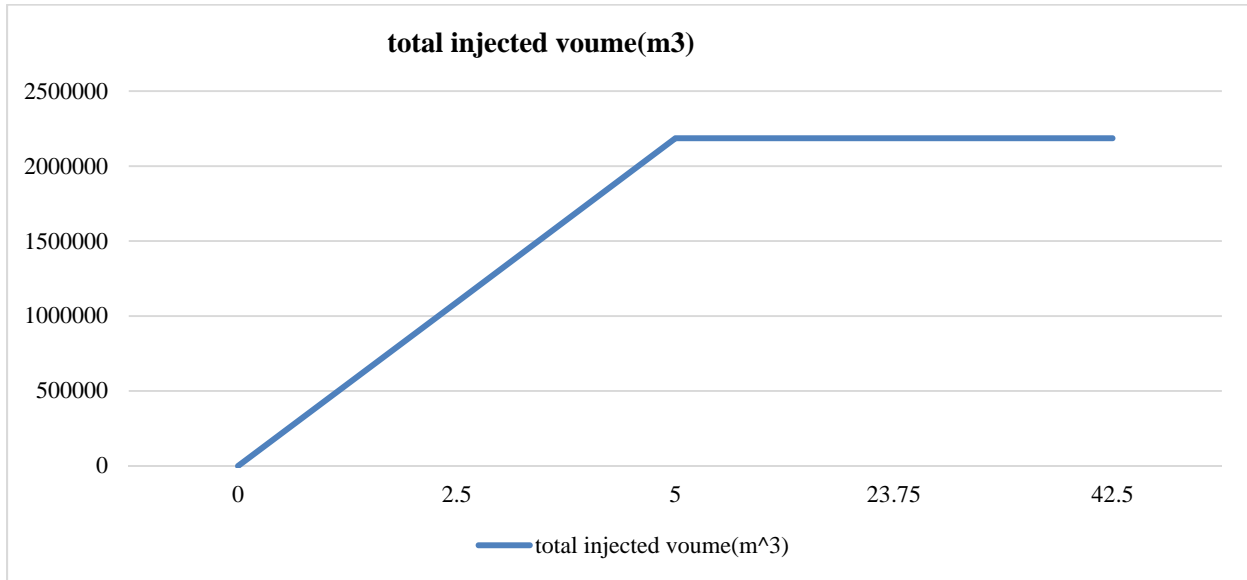


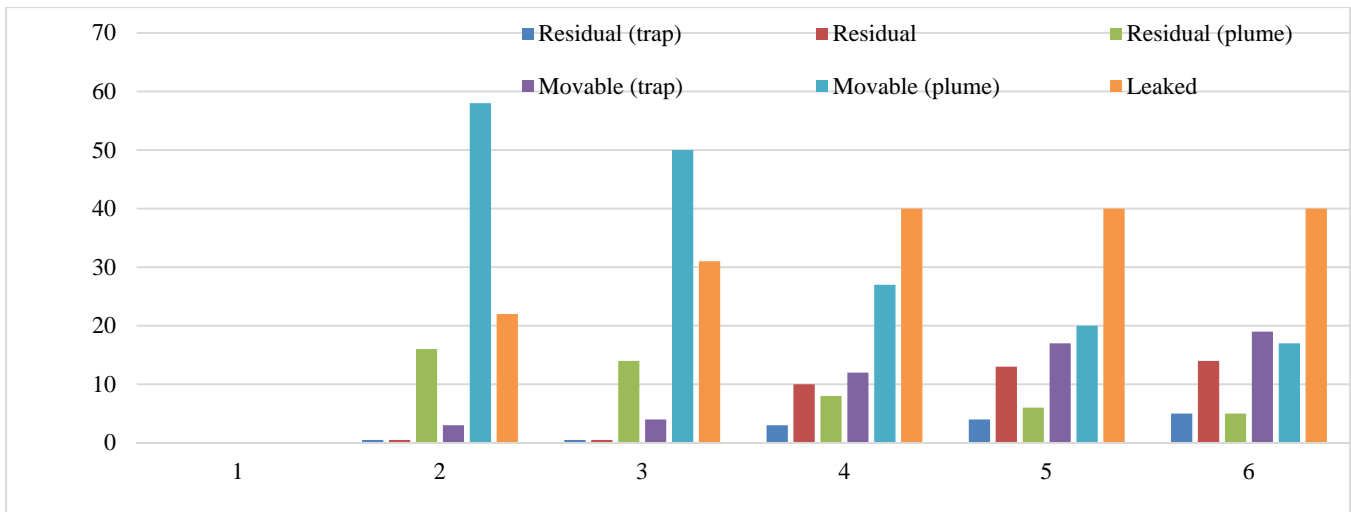
Fig. 5: The result of simulation at the end of migration. Here the Fig. illustrates the flow of CO₂ saturation at 80 years and due the buoyant effect the CO₂ is moving towards the cap rock

5.1. Graphs

The graphs are based on the results obtained at the time of simulation



The above graph will give the information about total injected volume at particular year (x-axis : years and y-axis: total injected volume)



The above graph is exposing the result of simulation at particular time steps discussing about residual and movable trap mechanism. Here x-axis: time steps and y-axis: percentage.

6. CONCLUSION

We have exposed the results of full 3D vertical equilibrium simulations of effective migration of CO₂ in a model of a site where CO₂ can be stored. The investigation of these results proves that for the specific case of CO₂ migration in the low permeable geological formations. Vertical equilibrium models present solutions that are more exact. The Vertical equilibrium model is much faster than equivalent 3D simulations that

resolve the similar dynamics. This is mainly important for simulations of long period migration of CO₂, wherever the plume thins out as it moves beyond from the injection site [I. Ligaarden et al (2010)]. Which create them more appropriate for sequential splitting approach. Based upon the results argued above, we propose that additional effort should go in to the development of more exact and faster vertical equilibrium models for simulating CO₂ migration. In addition, we look forward to that the renewed interest in vertical equilibrium models will result in the growth of new simulators that would be able to represent more complicated physical mechanisms that affect CO₂ migration in low permeable reservoirs.

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